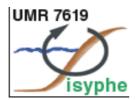
#### **ENGINE** mid-term Conference

### Recent progress in Discrete Fracture Network modelling for EGS development in tight fractured rocks

#### D. Bruel

Ecole des Mines de Paris, Centre de Géosciences, Systèmes Hydrologiques et Réservoirs UMR Sisyphe 7619, Université Pierre et Marie Curie





### Rationales for choosing Discrete Fracture Networks models Our experience comes from: (1/2)

- Nuclear waste management studies in hard rocks
  - EC programs (FI4W-CT96-0033, EUC19134 EN)
  - Benchmark DECOVALEX (NIREX data base, SKI Report 98:39)
  - ANDRA (HAVL granite, report GRPF DBG 05.0001.B)
- Water resources in semi arid granite areas
  - CEFIPRA projects in alterated basements, NGRI institute, India.
- Liquid-gas storage /water curtain systems, (GEOSTOCK/Tafjord reservoir, ...)
- Early HDR modelling experiences in shallow systems
  - french project (Mayet de Montagne)
  - UK project (Rosemanowes quarry)
- Soultz project, intermediate 3 km deep reservoir
  - EC programs (JOR3-CT95-0054, CT98-0313, STREP FP6-EGS)
  - Partnership EDRF, BRGM, ADEME

#### Rationales for choosing Discrete Fracture Networks models Lessons learned (2/2)

- Fractures exist at any scale
  - Very dense populations can be observed along wells
  - They reflect tectonic history (a classification is possible)
- Few fracture control flow along bore holes
  - Transport is highly heterogeneous
- Rock matrix is of very little importance for flow
- Interactions with stress regime are high (module+orientation)
  - Impact on global flow properties can be low
  - High pressures can be accommodated
  - Shear mechanism is not well understood
  - Released seismic energy can be large
- Role of rock damaged/alterated zones not properly accounted for

## Main assumptions to simulate fracture networks in rock blocks(1/2)

- Fractures are planar, disc shaped, and assemble in 3D networks
  - Density (from scan lines)
  - Size distribution
    - lognormal
    - Power law  $(R_0, a)$
  - No matrix
- Flow is not evenly 2D distributed in plane but concentrated in 1D channels
  - Fracture aperture
    - Cubic law:  $T \sim e^3$
    - Porous material:  $T \sim e$



## Main Assumptions Possible couplings on each fracture (2/2)

- Normal compliance versus normal effective stress
  - Normal closure law
- Shear rupture + normal dilation versus shear stress
  - Mohr-Coulomb criterion pre/post rupture friction, cohesion
- Thermal exchange conduction across adjacent rock blocks
  - Fluid density
  - Fluid viscosity
  - Induced thermo-elastic stress components
- Possibly two non-miscible fluids
  - Natural heavy brine fresh water

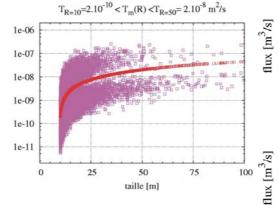
### Example of the DFN application in very tight systems (Rad Waste Management)

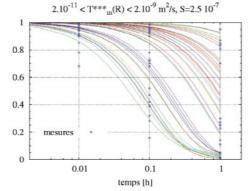
Many local flow experiment are available – Upscaling?

Straddle packer investigations in two wells (Charoux Civray site, ANDRA)

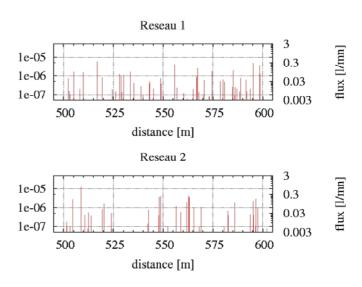
- Coupe Est-Ouest(km) 451.6 451.7 451.8 451.9 452.0 452.1 100 100 T [m<sup>2</sup>/s) 200 300 300 400 400 500 500 600 600 CHA112 700 700 800 800 P/P<sub>0</sub> [-] CHA212 900 900

Slug test analysis. A bimodal size-transmissivity distribution seems appropriate)



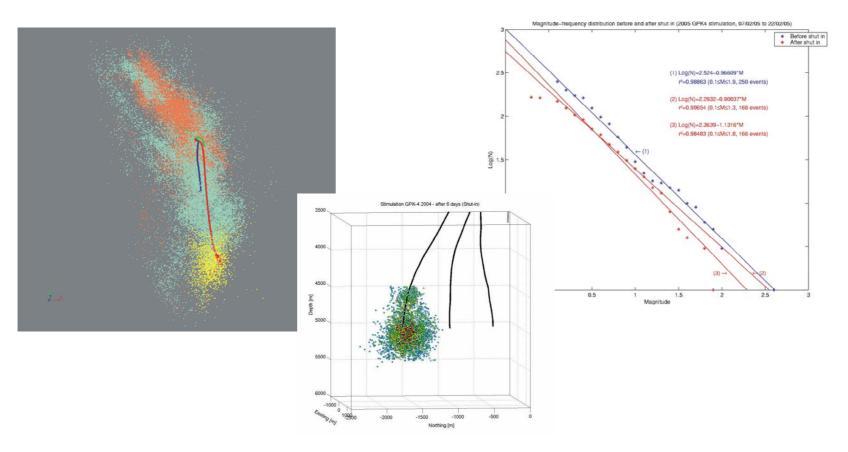


Upscaling: Flow log simulation at intermediate scale



Conclusion: 80% of the joints are partly sealed. Probable correlation between infillings and orientation. Non direct correlation between fract -ure size and hydraulic aperture.

# EGS studies (Soultz sous Forêts site) Local hydraulic tests are sparse (hot/expensive/risky) but Some large scale indirect characterization is available

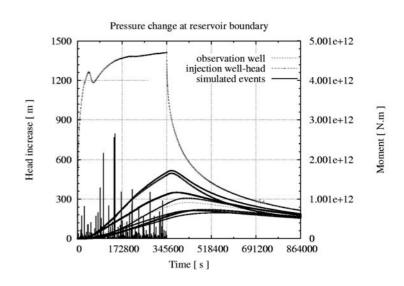


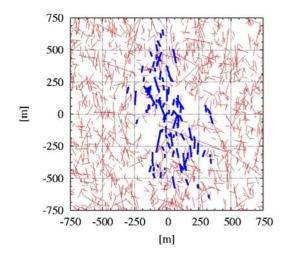
3D DFN geometry and initial transmissivity can be constrained by the observed microseismic migration. Magnitude frequency diagrams gives information on the size distribution of the fractures. Identified large structures can be superimposed

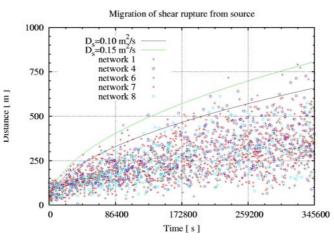
### Discussion of network properties against seismic

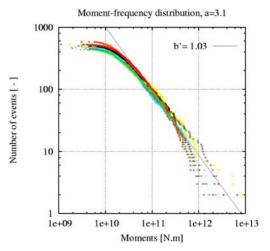
data (after Bruel, in *Int. J. Rock Mech. & Abstr.* In press 2007)

3D Random networks are generated and hydraulic tests at high rates are performed (GPK4 stimulation in 2004 as an example) Shear rupturing processes are simulated and analysed as a diffusion process. Hydraulic diffusivity is in the order of  $0.15 \text{ m}^2/\text{s}$  (K  $\sim 1.10^{-17} \text{ m/s}$ ) and a power law exponent for the fracture size is proposed. (a=2.7)



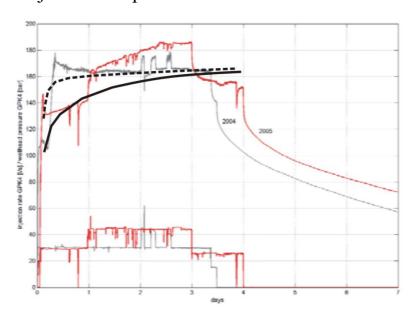


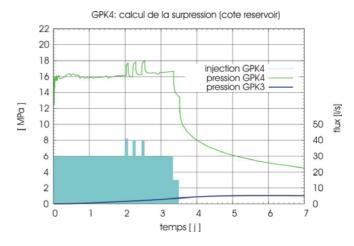


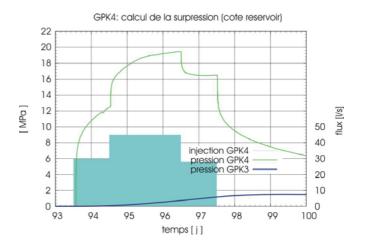


# Discussion on hydraulic efficiency of hydraulic stimulations against GPK4 measurements (2004,2005) (1/2)

GPK4 was stimulated in 2004 and again in 2005 (~30000 m³ each). No structure was developed at large scale. The model is used to understand the increase of hydraulic capacity of the fracture network around the well. The role of a nearby structure delineated as a 'no seismic zone 'can be discussed. This zone would separate the reservoir into two adjacent compartments.





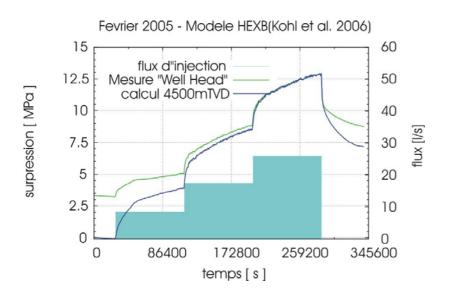


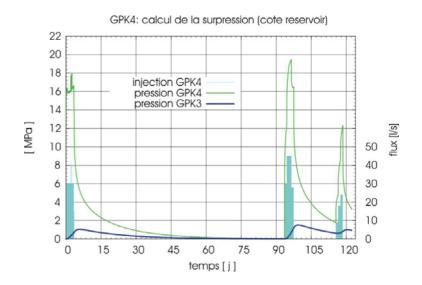
## Discussion on hydraulic efficiency of hydraulic stimulations against GPK4 measurements (2004,2005)

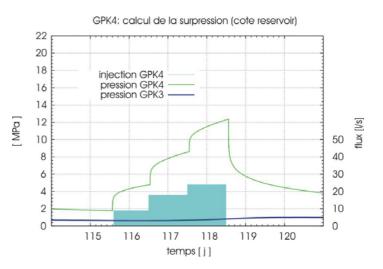
(2/2)

Post stimulation tests were performed to evaluate the impact on fluid injectivity. (Below a step rate test, involving  $\sim 10000 \text{ m}^3$  of water).

The model used in step 1 is run for the entire period. Step rates responses are well reproduced. However no steady state can be obtained within 3 days and no solid conclusions can be drawn from this test, regarding the gain in hydraulic injectivity of this well.





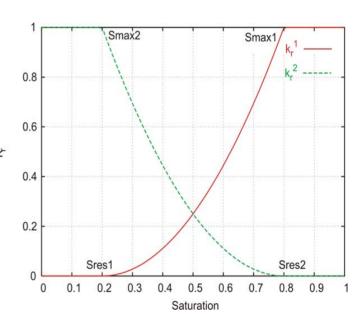


### Importance of alterations within the fractures. Role of fracture porosity in mass transfer: a two phase flow approach (1/3)

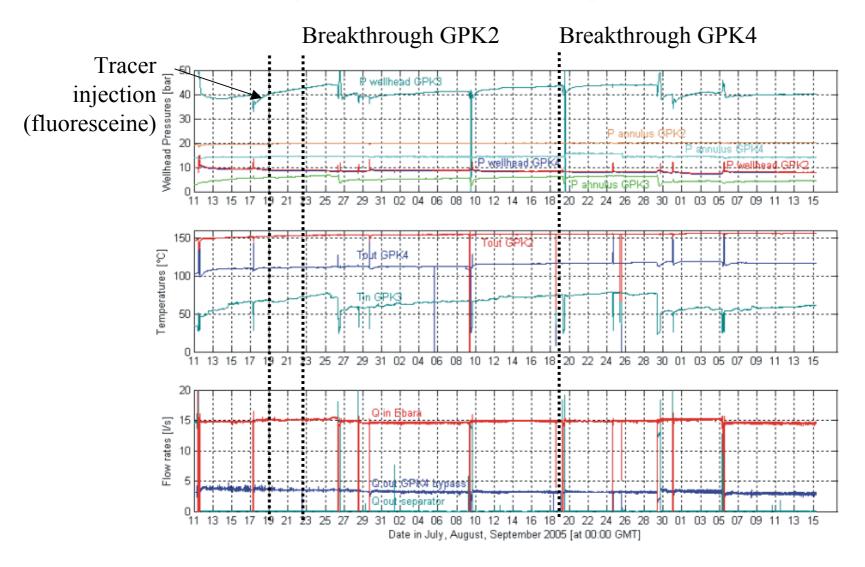
• Pressure /saturation formulation, for each fluid, assumed non miscible

$$\begin{cases}
\sum_{j} k_{ij} \frac{k_{ij}^{l1}}{\mu_{ij}^{l1} L_{ij}} \left( \left( P_{i}^{l1} - P_{j}^{l1} \right) + \left( \rho_{i}^{l1} g z_{i} - \rho_{j}^{l1} g z_{j} \right) \right) = S_{i}^{l1} C_{i} \frac{\partial P_{i}^{l1}}{\partial t} + \Phi_{i} \frac{\partial S_{i}^{l1}}{\partial t} \\
\sum_{j} k_{ij} \frac{k_{ij}^{l2}}{\mu_{ij}^{l2} L_{ij}} \left( \left( P_{i}^{l2} - P_{j}^{l2} \right) + \left( \rho_{i}^{l2} g z_{i} - \rho_{j}^{l2} g z_{j} \right) \right) = S_{i}^{l2} C_{i} \frac{\partial P_{i}^{l2}}{\partial t} + \Phi_{i} \frac{\partial S_{i}^{l2}}{\partial t}
\end{cases}$$

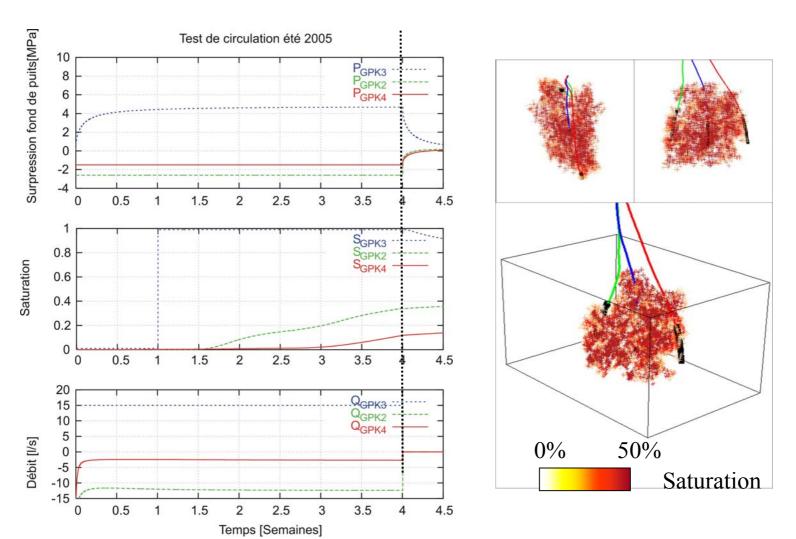
- $\rho$ ,  $\mu$  fluid density and viscosity
- S<sub>0</sub> initial saturation in in-situ fluid
- $\Phi$  fracture porosity  $\rightarrow$  volume for fluid storage is each fracture
- k<sup>l</sup><sub>ij</sub> relative fluid perméability
- IMPES numerical scheme+ Newton Raphson iterates on pressure



# Using the two phase flow module to discuss fracture volume. Application to the circulation test (Aug. Oct. 2005) (2/3) After Baujard et al. *Geothermics*, In press, 2007



# Using the two phase flow module to discuss fracture volume. Application to the circulation test (Aug. Oct. 2005) (3/3) After Baujard et al. *Geothermics*, In press, 2007



#### Discussion and Conclusion

- Discrete Fracture Network is a valuable approach because :
  - It captures heterogeneity of structure and hydraulic responses
    - Open thin fractures and planar porous zones can be mixed as 'objects'
  - It allows some non-reversible interactions between hydraulic and mechanical parameters
  - It provides a basis for a first interpretation of seismic activity
    - Delayed activity is obtained as a result of low diffusivity
    - Fracture size and large events could be linked
  - It is appropriate for mass transfer predictions and long term thermal calculations
- Progress are still required, among them
  - The understanding of coupled processes during shear motion
  - The response of rock damaged/alterated zones, in conjunction with thermal exchange and acidization experiments